

Spatial Variability of Chosen Soil Properties on Mountainous Area on Example of the Kasińczanka Stream Basin, Western Carpathians

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ABSTRACT

The phenomenon of erosion on mountain and submontane areas influences directly on high variability of soil properties. In the work there were presented results of analysis of spatial variability of bulk density, total porosity, organic matter content and saturated hydraulic conductivity, on eroded slopes of the Kasińczanka stream basin. Geostatistical analysis was carried out using the kriging method, based on irregular network, consisted of 52 points, situated by means of the GPS. Taking into account the calculated variability coefficient it was stated, that on the investigated area, the most flexible spatially was saturated hydraulic conductivity, while the less flexible turned out total porosity. Using the determined models of semivariance, the maps of spatial variability of chosen parameters were drafted. It was stated that high value of variance influenced on higher smoothing of spatial distribution in interpolation. Results of geostatistical analysis will allow to find locations for new measuring points, what has substantial significance in mountain areas, for precision analysis of soil properties. Based on the obtained results, it can be stated that the kriging method may be useful tool for determination spatial variability analysis of soil properties on an areas of mountain basins.

Keywords: spatial variability, semivariogram, kriging, geostatistics

INTRODUCTION

The physical, water and chemical properties of both soils and waste are characterized by spatial variability [Boroń et al. 2016, Duffera et al. 2007, Klatka et al. 2016, Medina et al. 2017, Najafian et al. 2012, Tola et al. 2017, Xiong et al. 2015]. The anthropogenic activity influences this variability [Barik et al. 2014, Bogunovic et al. 2017, Klatka et al. 2015, Rosemary et al. 2017, Usowicz and Lipiec 2017, Wei et al. 2009]. Particularly, differentiation of soil properties occurs on slope areas as a result of water erosion [Halecki et al. 2017]. This problem, occurring on lowland areas, was investigated by many authors [Rejman et al. 1998, Williams et al. 1984, Stach 1998]. However, such investigations on areas of mountainous basins were carried out on a small scale. The direct consequence of the evaluation pertaining to the spatial variability of soil environment proper-

ties may be the lack of agrotechnical measures, which are not adapted to the predominant conditions [Usowicz et al. 2004]. The knowledge on the spatial variability of soil properties in mountainous stream basins has the basic significance in the determination of the actual conditions of plant growth and development. It is influenced by the phenomena resulting from the processes having relationship with topography and land use [Pennock and De Jong 1987]. Under the mountainous conditions, the land use is subjected to a high degree to configuration and climatic as well as soil conditions. As a result of erosion processes, the agricultural fields on such areas are characterized by high spatial variability of soil properties, which is influenced by numerous factors. The real variability is virtually never linear. That is why there is a need for using interpolation methods, which give better possibility of recognizing the soil properties regarding the specified charac-

ter of spatial variability. One of such methods is kriging, which allows to determine the mean value of a given variable property in any part of the considered area and to find a location for new measuring points, which will narrow the section of estimation to the highest degree. The spatial variability of soil properties is most often investigated based on the samples taken in regular or irregular measurements [Klatka 2009].

The aim of this work was to conduct the analysis of spatial variability of selected soil properties on the area of a mountain basin using the geostatistical kriging method.

MATERIALS AND METHODS

The investigations were carried out on the Kasińczanka stream basin, extending over the area of 48,54 km², located in the Limanowa administrative district, Mszana Dolna community (Small Poland voivodship). The main stream, with the length of 17,74 km, constitutes the inflow of the Raba river. The basin is closed by uplifts belonging to the Mesoregion of the Beskid Wyspowy – slopes of the mountains: Lubogoszcz (967 m a.s.l.), Śnieżnica (1005 m a.s.l.), Wierzbanowska (774 m a.s.l.), Dzielec (607 m a.s.l.) and slopes of Kiczara (710 m a.s.l.). The study work was based on the cartography analysis of the topographic map 1:5000, Digital Elevation Model (DEM) of raster resolution 5 m, and orthophotomap 1:10000, using the MapInfo Professional 15.0 [MapInfo 2002] and the Surfer 8 [Golden 2002] software.

Within the framework of the field investigations on the area of the Kasińczanka stream basin, 52 samples of soil were taken from the upper layer (0–10 cm) in three replications. The location of points is presented in Figure 1b. In the course of laboratory investigations, the following soil properties were determined: bulk density, total porosity, organic matter content and saturated hydraulic conductivity. The physical properties (bulk density and total porosity) were determined using the rings of 100 cm³ volume [Mocek et al. 1997]. Saturated hydraulic conductivity was determined in laboratory, using the method of hydraulic pressure fall, in the apparatus based on the Darcy law, with controlled pressure height and electronic reading of water volume. Organic matter content was determined using the Turin method, modified by Oleksynowa et al. [1991],

consisting in the oxidation of humus by potassium bichromate (Cr⁶⁺).

The following statistical measures were calculated with the aim of characterizing the investigated lithogenic parameters and carrying out preliminary determination of their variability degree [Hellwig 1993]: maximum value, minimum value, median, mean value, variance, standard deviation and variation coefficient. Variation coefficient was determined as measure of variance scattering and standard deviation in a shape of a given characteristic quotient – standard deviation and mean value of the given characteristic, expressed in percent. The results were presented in Table 1. Spatial variability of the investigated soil properties was determined using the kriging method. This method allows to estimate the confidence intervals of estimation, determine the mean value in any part of an investigated area and to find location for new measuring points, which to the highest degree will narrow measured in point intervals of distributions. Specifying the term kriging, it should be stated that this is method optimizing estimation of spatially correlated the quantity Z , both in the case of its stationarity and for nonstationarity instances. If we assume that $Z_i = Z(x_i)$ is a quantity measured in point x_i ($i = 1, 2, \dots, n$), the problem of point estimation consists in determination of value Z_0 in point x_0 , where value Z was not measured. By continuous change of location of point x_0 it is possible to determine all the field of variation of variable Z . The analysis was conducted in two stages. In the first one, the correlation of distribution in a shape of the function $\gamma(h)$ stated in geostatistics by empirical semivariogram was determined [Smith et al. 1994]. The obtained results allowed to carry out the quality description of regularities occurring in the differentiation of values of the analyzed parameters. The empirical semivariogram course was determined in the shape of points calculated according to the equation [Oliver and Webster 1986]:

$$\gamma(h) = \frac{1}{2m(h)} \cdot \sum_{i=1}^{m(h)} [Z(x_i) - Z(x_i + h)]^2 \quad (1)$$

where: $\gamma(h)$ – value of semivariogram for a distance h ,

$m(h)$ – number of observations outlying of distance h ,

$Z(x_i)$ – value observed in point x_i ,

$Z(x_i + h)$ – observed value in point outlying of distance h .

The determined points of empirical semivariograms were described by means of mathematical functions and the theoretical semivariogram was elaborated. The positively determined functions, with zero value in point zero were used for their description. Marx and Thompson [1987] presented the following models: spherical, exponential, Gauss', linear and logarithmic; however, the functions belonging to so called safety class – linear and spherical functions, are used most often. The obtained shape of semivariogram points the kind of spatial correlation of the considered variable. There are two substantial kinds of semivariograms. The first one is characterized by the fact that the variance value increases when the distance h rises up to a maximum value and remains in this state together with the increase of a distance. After the initial increase, value $\gamma(h)$ attains for a given distance α the value equal to variance of a variable. Distance α is called the extent of semivariogram and determines the limit of spatial correlation. Variable Z , having semivariogram of this kind, is stationary not only internally, but also has stationarity of the second order. In the second type of semivariograms, variance $\gamma(h)$ is increasing function without limitations. In this case, variance is unlimited and with an increase of distance h , attains some boundary value, denoted as C . The boundary value is attained exactly or only asymptotically. When semivariogram does not pass the beginning of the coordinate system, then this is called nugget effect, specified as c_0 . This effect may testify that in the case of taking nugget to a sample, other samples – even being very close – will differ by the degree of concentration characterizing the given component.

Maps of spatial variability of the investigated soils were determined using digital procedures – Surfer 8 [Golden 2002]. The input data for this program were determined by means of the STATANAL software [Stein and Staritsky 1993], which allows spatial interpolation on the basis of functions of semivariograms and determined parameters of smoothed models (nugget effect, sill effect), and character of variables. Cutting of isolines was adjusted to the quantity of random error – nugget variance.

RESULTS AND DISCUSSION

From the point of view of land use and cover, cultural agriculture-forest landscape prevails in the investigated area (Figure 1a). High degree of

forestation (50,19%) is characteristic to this area. There are mainly coniferous forests, with predominant tree species: *Picea abies*, *Abies alba*, *Pinus sylvestris*, *Fagus sylvatica*, *Betula pendula* and *Alnus incana*. The calculated index of forestation development is equal to 0.487 and is not advantageous for the flows of water in the investigated basin [Lipski et al. 1980]. The area of agricultural production space is used intensively. Small individual farms prevail. Arable lands occupy 63 km², which corresponds to 42.50% of the total area. In the structure of land use in individual farms, grasslands (23.27%) and arable lands (13.78%) prevail. The remaining 7.31% of the area is occupied by rural development and traffic grounds. The density of hydrographic network amounts to 2.21 km×km⁻² (Figure 1d). Mean weighted slope for the entire basin is 16.12%. Soils in the basin were created as a result of weathering of sandstones, shales and marls. Shallow soils, skeletal in the form of light sandy-clay loams prevail. In the higher parts of the basin, on more steep slopes, loamy soils occur, in the form of sandy-clay and sandy loams and clayey silt. In the south-west part of the basin, on lower parts sandy soils in the form of loamy sands, silty and clay silt soils are found. These soils occupy highly eroded places, creating ravines of streams and deeply cut roads, partly covered by grasses and shrubs, not forested because of the highly intensive denudation processes (Figure 1c).

The determined values of bulk density fluctuated between 0.75 and 1.53 Mg·m⁻³, attaining the mean of 1.31 Mg·m⁻³. The values below 1.25 Mg·m⁻³ were determined in the samples taken from forest area. It can be stated that bulk density was influenced by land use. Total porosity ranged between 40.2 and 67.4%, 48.48% on average. Because total porosity influences, among others, the air properties and water permeability, it can be stated that in the case of most points the values are advantageous regarding the resistance to erosion. The content of organic matter in the investigated samples is within the range of 1.85–4.28%. These are typical values for the given subtypes of soils in mountain areas [Mocek 2015]. Regarding this property, the investigated soils are not exposed to erosion. Saturated hydraulic conductivity ranged between $0.84 \times 10^{-3} \text{ m} \cdot \text{d}^{-1}$ and $4.87 \times 10^{-1} \text{ m} \cdot \text{d}^{-1}$. Water permeability plays a great role in surface runoff forming. High values of this parameter ($\geq 8.64 \text{ m} \cdot \text{d}^{-1}$) show low surface runoff; on the other hand, the lower ones ($< 0.000864 \text{ m} \cdot \text{d}^{-1}$) cause a higher

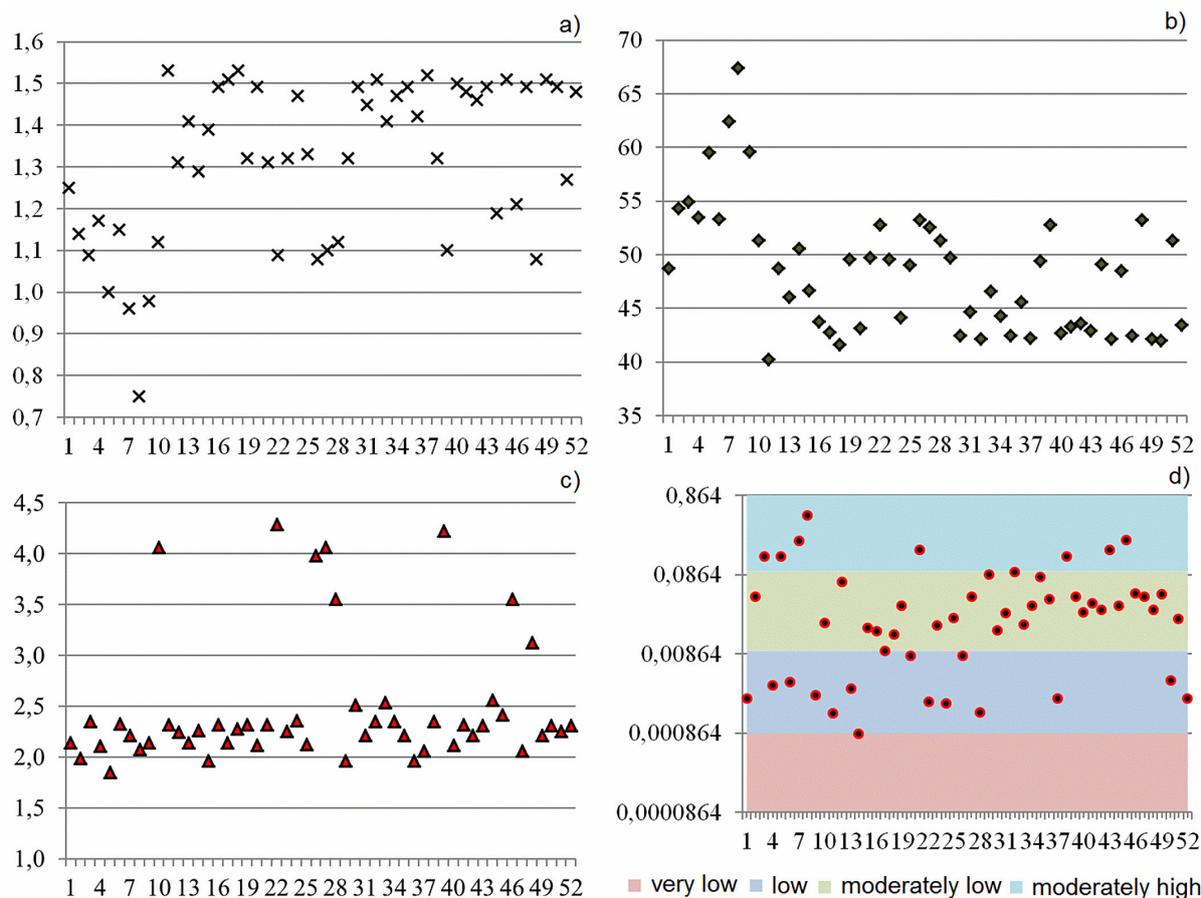


Figure 2. Results of parameters determination: a) bulk density ($\text{Mg}\cdot\text{m}^{-3}$), b) total porosity (%), c) organic matter content (%), d) saturated conductivity coefficient ($\text{m}\cdot\text{d}^{-1}$)

Table 1. The determined statistical measures

Investigated parameter	Statistical measures					
	Min value X_{min}	Max value X_{max}	Mean value X	Variance $V(X)$	Standard deviation $\delta(X)$	Variability coefficient $\beta[\%]$
Bulk density ($\text{Mg}\cdot\text{m}^{-3}$)	0.75	1.53	1.31	0.04	0.26	19.92
Total porosity (%)	40.20	67.40	48.48	40.07	6.37	13.14
Organic matter content (%)	1.85	4.28	2.49	0.43	0.71	28.6
Saturated conductivity coefficient ($\text{m}\cdot\text{d}^{-1}$)	0.00084	0.4780	0.0618	0.0098	0.0561	90.78

value of variation influenced the greater smoothness of spatial distribution in interpolation using the kriging method. Similar relationships were stated in the investigations carried out by Klatk et al. [2017].

Spatial variability of mineral soils properties using the geostatistical kriging method was the subject of investigations carried out by many authors. Wei et al. [2006] used this method for the evaluation of spatial variability of organic matter content in the north-east part of China. Shao et al. [2006] determined spatial variability

of nutrients distribution in soil. Stach [1998], in turn, determined the spatial distribution of arable layer properties on litologically heterogeneous morainal slope. The results of investigations of the afore-mentioned authors show that the geo-statistical analysis and maps of spatial variability enable the estimation of the investigated variable value in every point of the considered area and determination of mean value in any part of the area. It allows to find the location for new measuring points, which narrow the estimation intervals of distributions to the highest degree. Taking

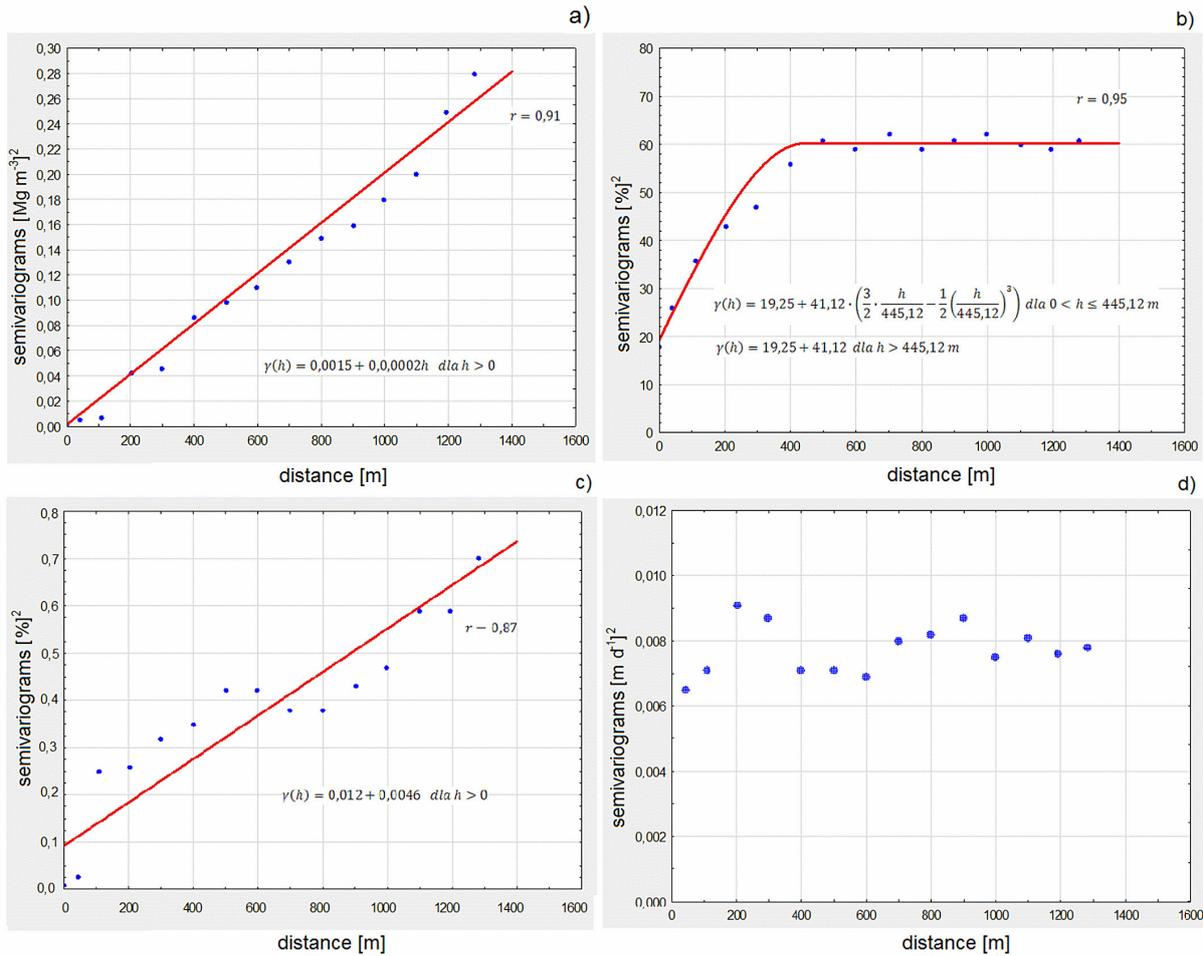


Figure 3. Empirical and theoretical semivariograms: a) bulk density ($\text{Mg} \cdot \text{m}^{-3}$), b) total porosity (%), c) organic matter content (%), d) saturated conductivity coefficient ($\text{m} \cdot \text{d}^{-1}$)

Table 2. Parameters of the fitted models

Soil properties	Model	Nugget effect. c_0	Boundary effect. C	Scope. α [m]	Correlation coefficient. r
Bulk density	linear	0.0015	0.0002	-	0.91
Total porosity	spherical	19.25	41.12	445.12	0.95
Organic matter content	linear	0.012	0.00046	-	0.87
Saturated conductivity	-	-	-	-	-

into account the erosion processes in mountain streams basins this problem has the basic significance for determination of the sampling distribution for precise analysis of soil properties.

The drafted spatial distribution maps of the investigated properties show their high relation with land use and runoff processes on slopes and the erosion connected with it. Moreover, spatial distribution of the investigated parameters show high dynamics, which – in the case of arable land – is influenced by time that passed by from last tillage treatment.

CONCLUSIONS

Regarding texture, bulk density and total porosity, the soil on the investigated area is relatively resistant to water erosion. However, low values of saturated hydraulic conductivity point to the possibility of high runoff occurrence. Spatial variability of the investigated soil properties is influenced to a high degree by land use.

The theoretical semivariogram for saturated hydraulic conductivity and limit of spatial correlation were not determined for the investigated

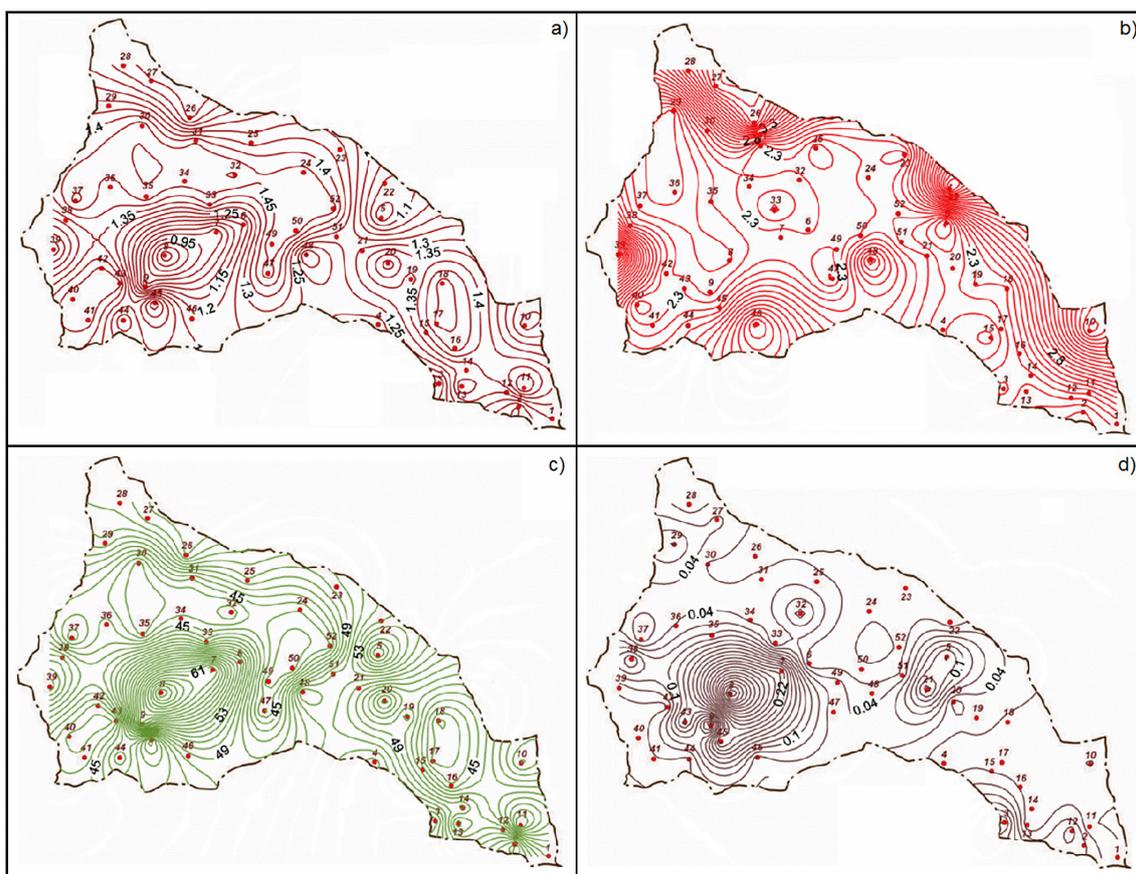


Figure 4. Spatial distribution on the basin area: a) bulk density ($\text{Mg}\cdot\text{m}^{-3}$), b) total porosity (%), c) organic matter content (%), d) saturated conductivity coefficient ($\text{m}\cdot\text{d}^{-1}$)

soils. This is connected with high random variability of these properties. For bulk density and organic matter content, semivariation increased linearly with distance, and it was not possible to determine the limit of spatial correlation as well. In the case of total porosity, the limits of spatial correlation were determined. The nugget effect was observed which may testify the occurrence of variability of a given parameter in a distance lesser than the one between measuring points.

The stated spatial relationship of the investigated soil parameters shows that for their representative description, the statistical evaluation is not sufficient, because of the basic assumption of statistics – independent of samples. Because classical statistics are reliable only beyond the area of spatial dependence that is why significant error may occur within it. The determined relationships of the investigated properties show that during the calculation of basic statistics for a given area, the character of their relations has to be taken into consideration. On the basis of the obtained results, it can be stated that the kriging method may be a useful tool for determination spatial distribution of soil properties in mountainous area.

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